

Balloon Data and Planetary Temperature Profiles

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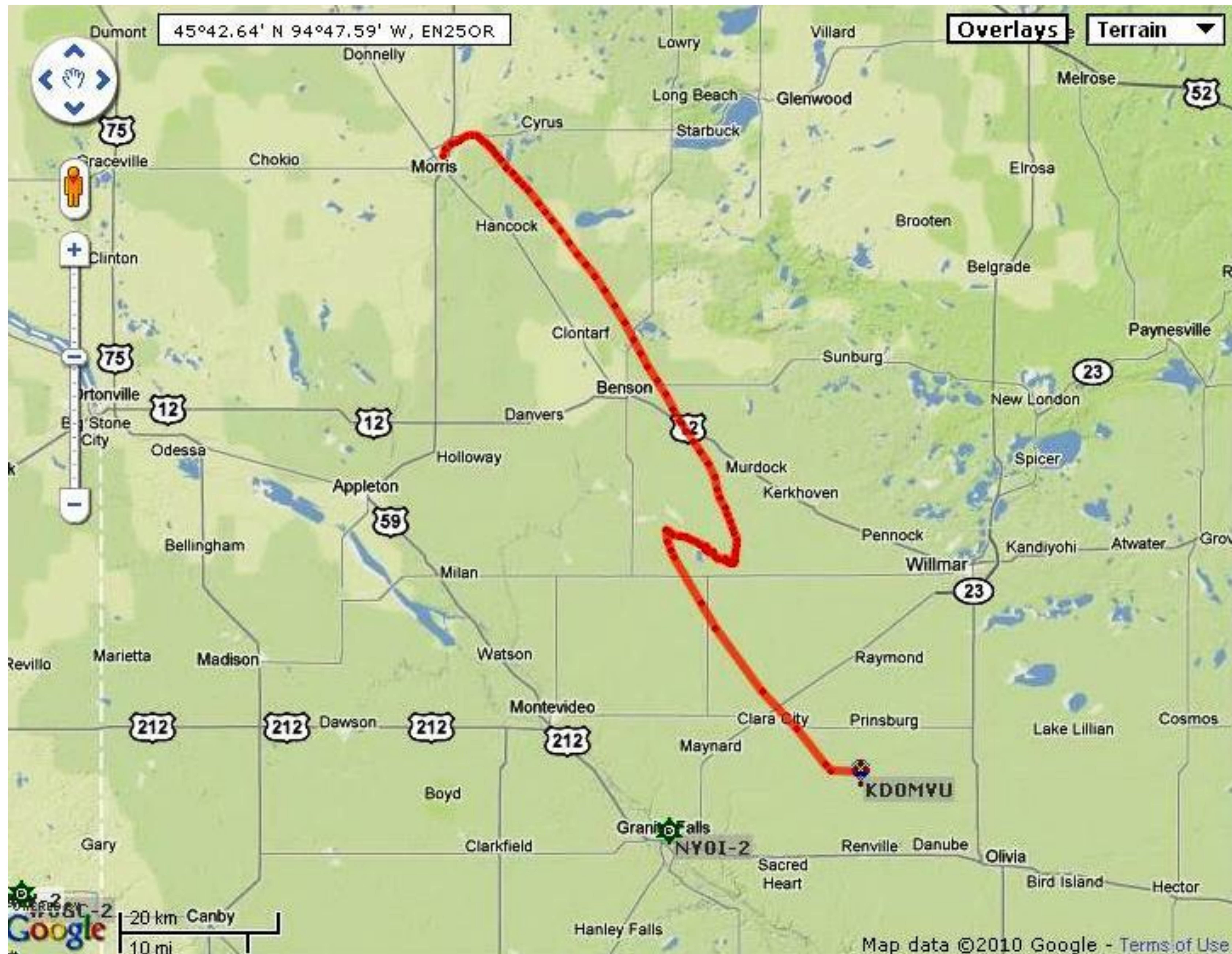


Outline

- Launch information
- Temperature measurement
- Earth's atmosphere
- Atmospheric stability
- Solar system bodies with solid surfaces and substantial atmospheres
- Lab ideas



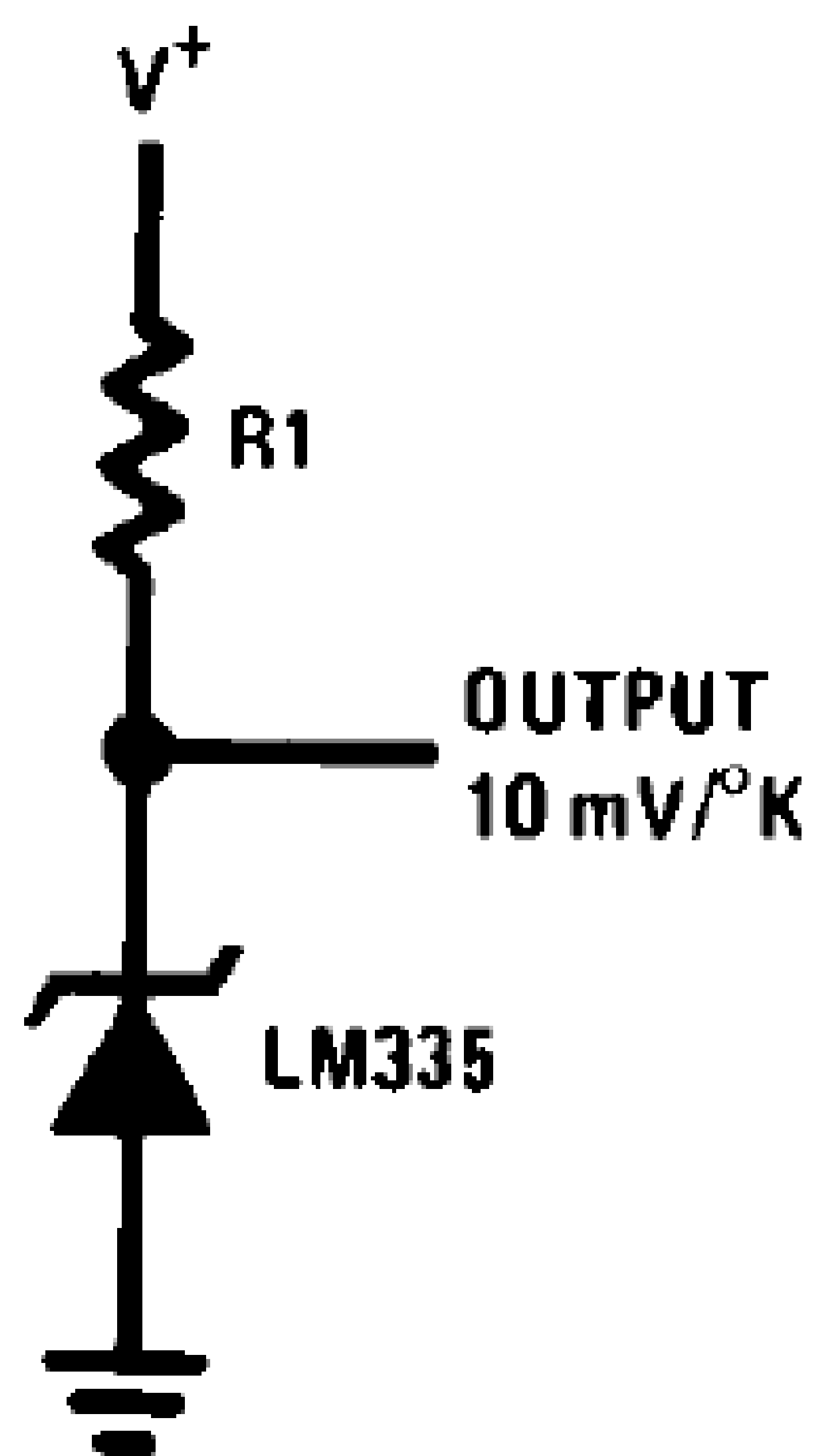




Temperature Circuit

R1 2K Ω

Basic Temperature Sensor



TL/H/5698-2

DeviceId	Timestamp	A1	A2	A3	A4	A5	A18
30	11/6/2010 8:52	553	874	571	585	855	351
30	11/6/2010 8:52	540	867	573	584	851	412
30	11/6/2010 8:52	534	865	573	585	849	451
30	11/6/2010 8:52	503	856	574	585	854	552
30	11/6/2010 8:52	497	853	574	586	855	588
30	11/6/2010 8:53	476	843	576	585	850	684
30	11/6/2010 8:53	473	841	577	586	849	715
30	11/6/2010 8:53	451	835	579	586	850	783
30	11/6/2010 8:53	434	832	580	587	854	809

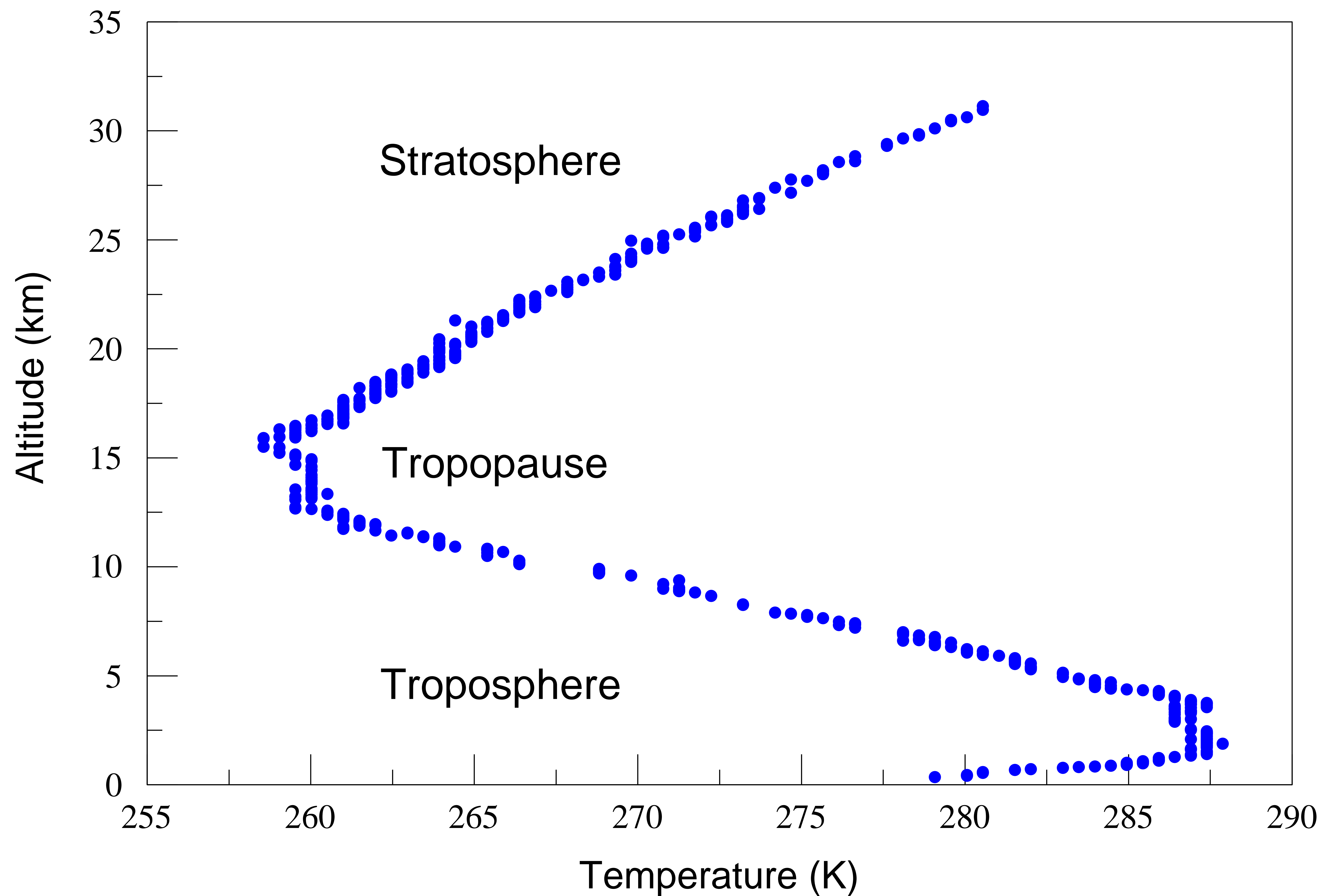
A18 Altitude from GPS position

A3 Digitized External Temperature

$$T(K) = T(V) * 5/1023 * 100$$

Altitude versus Temperature

6 November 2010



Troposphere

normally heated from the bottom

near surface - inversion

8:52 am, 6 November 10

measured environmental

lapse rate (ELR) 3.2 K km^{-1}

average ELR

$\sim 6.4 \text{ K km}^{-1}$

Tropopause

lapse rate $\sim 0 \text{ K km}^{-1}$

Stratosphere

temperature gradient $+1.5 \text{ K km}^{-1}$

heating due to solar ultraviolet radiation

photo-dissociation of O_3

Atmospheric Stability in the Troposphere

Rising air bubbles cool.

If the bubble cools faster than the surrounding air it will sink back down and be stable

If the bubble cools more slowly than the surrounding air it will continue to rise.

These temperature changes are indicated by the lapse rates.

Environmental lapse rate (ELR)

measured $|\Delta T/\Delta z|$

Dry adiabatic lapse rate (DALR)

9.76 K km^{-1}

Saturated adiabatic lapse rate (SALR)

minimum $\sim 4 \text{ K km}^{-1}$

DALR Derivation

First Law

$$dQ = dU + \delta W$$

of

$$= nc_v dT + PdV$$

Thermodynamics

Ideal Gas Law

$$pV = nRT$$

$$Vdp + pdV = nRdT$$

$$R = c_p - c_v$$

$$dQ = nc_v dT + nRdT - Vdp$$

$$= nc_p dT - Vdp$$

$$\text{Adiabatic (dQ = 0)} \quad nc_p dT = V dp$$

Molecular mass per mole = m_m

$$C_p = c_p / m_m$$

$$\rho = n m_m / V$$

$$\begin{aligned} dT / dp &= V m_m / (n c_p m_m) \\ &= 1 / (C_p \rho) \end{aligned}$$

Hydrostatic
Equilibrium

$$dp = -g\rho dz$$

$$dT/(-g\rho dz) = 1/(C_p \rho)$$

so

The DALR

$$dT/dz = |-g/ C_p|$$

$$g_{\oplus} = 9.80 \text{ m s}^{-2}$$

$$C_p_{\oplus} = 1004 \text{ J Kg}^{-1} \text{ K}^{-1}$$

$$\text{DALR}_{\oplus} = 9.76 \text{ K km}^{-1}$$

SALR lower due to the release of the
latent heat of condensation
(minimum $\sim 4 \text{ K km}^{-1}$)

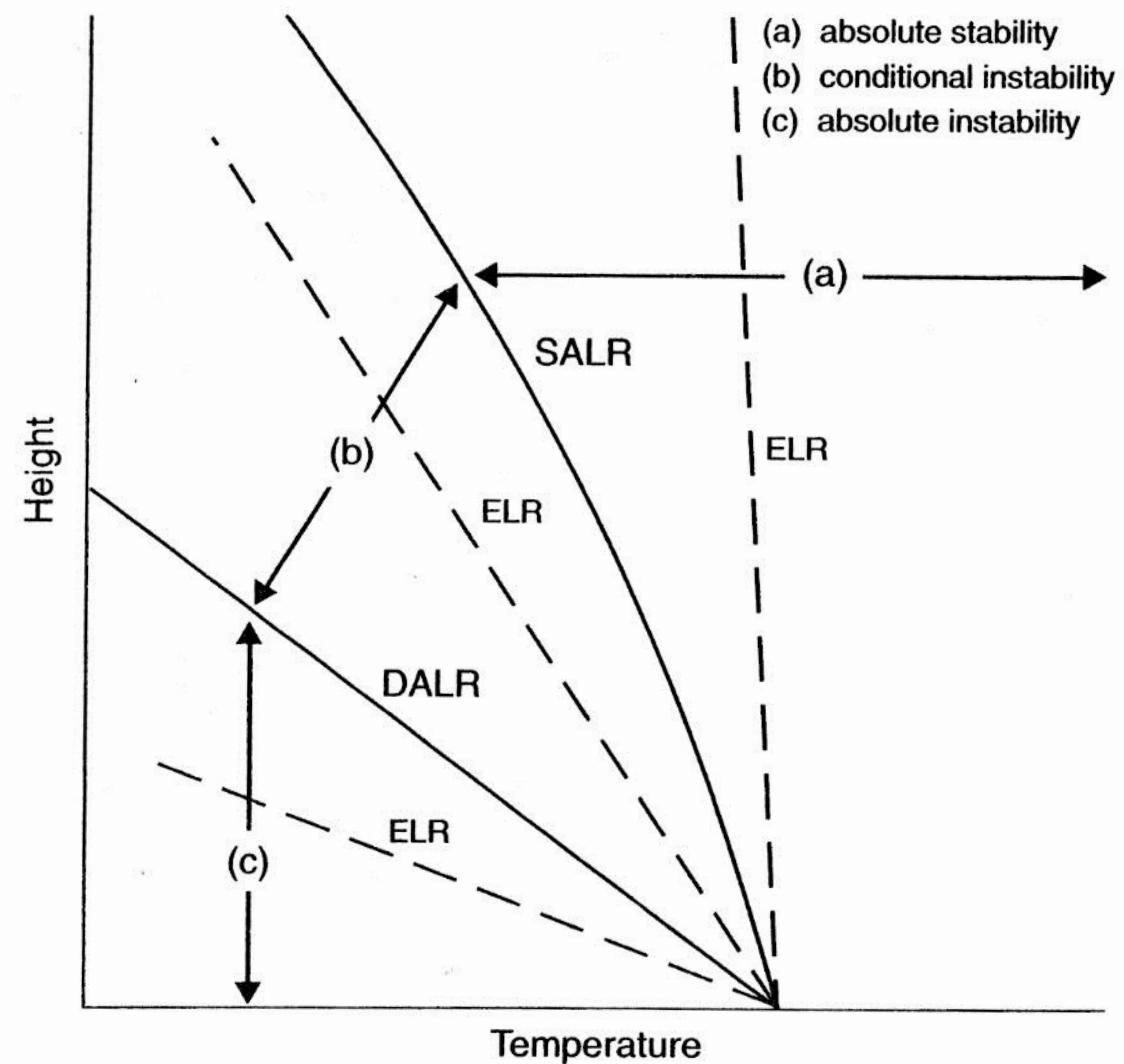
Stability

Absolute stability
 $ELR < SALR$

Conditional instability
 $SALR < ELR < DALR$

Absolute instability
 $ELR > DALR$

6 Nov 10 launch
 $ELR = 3.2 \text{ K km}^{-1}$



Solar System Bodies with Solid Surfaces and Substantial Atmospheres

Atmospheric Constituents

Earth	78% N ₂	21% O ₂	
Venus	96.5% CO ₂	3.5% N ₂	
Mars	95.3% CO ₂	2.7% N ₂	1.6% Ar
Titan	90-97% N ₂	0-6% Ar	0.5-4% CH ₄

Atmospheric Pressures

Earth 1010 hPa

Venus 9.2 MPa

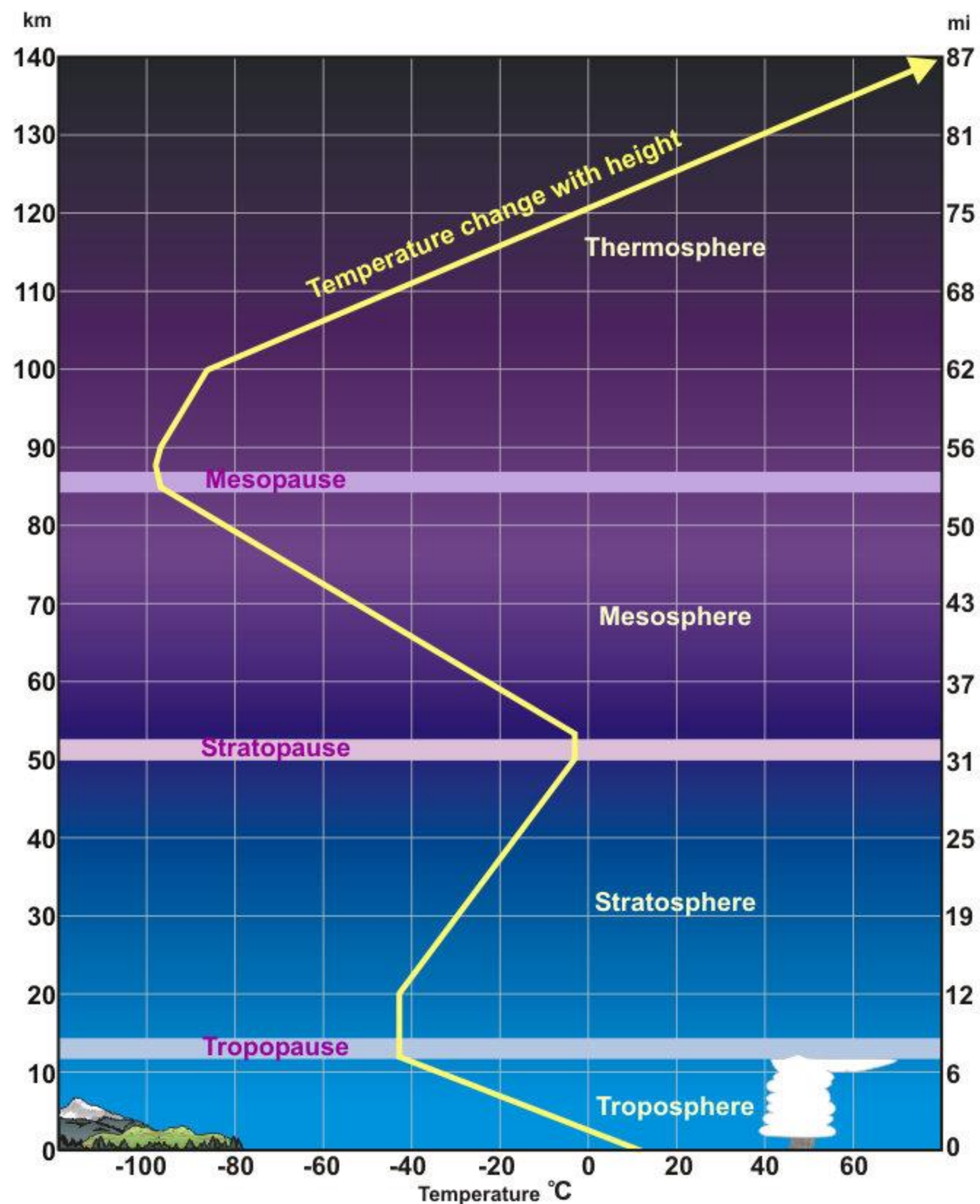
Mars 6 hPa

Titan 1500 hPa

Earth

Altitude versus Temperature

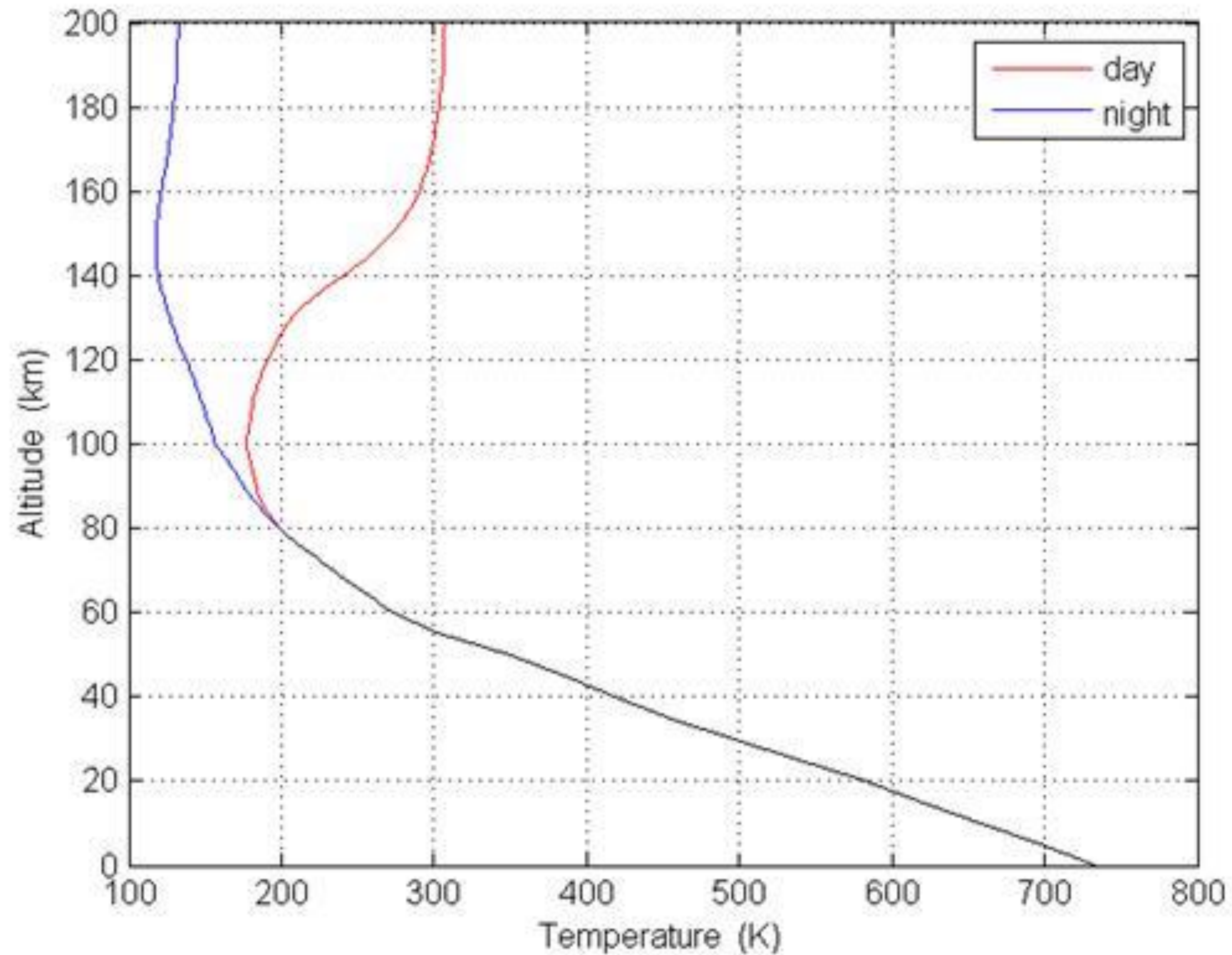
<http://www.srh.noaa.gov/srh/jetstream/atmos/atmprofile.htm>



Venus

Altitude versus Temperature

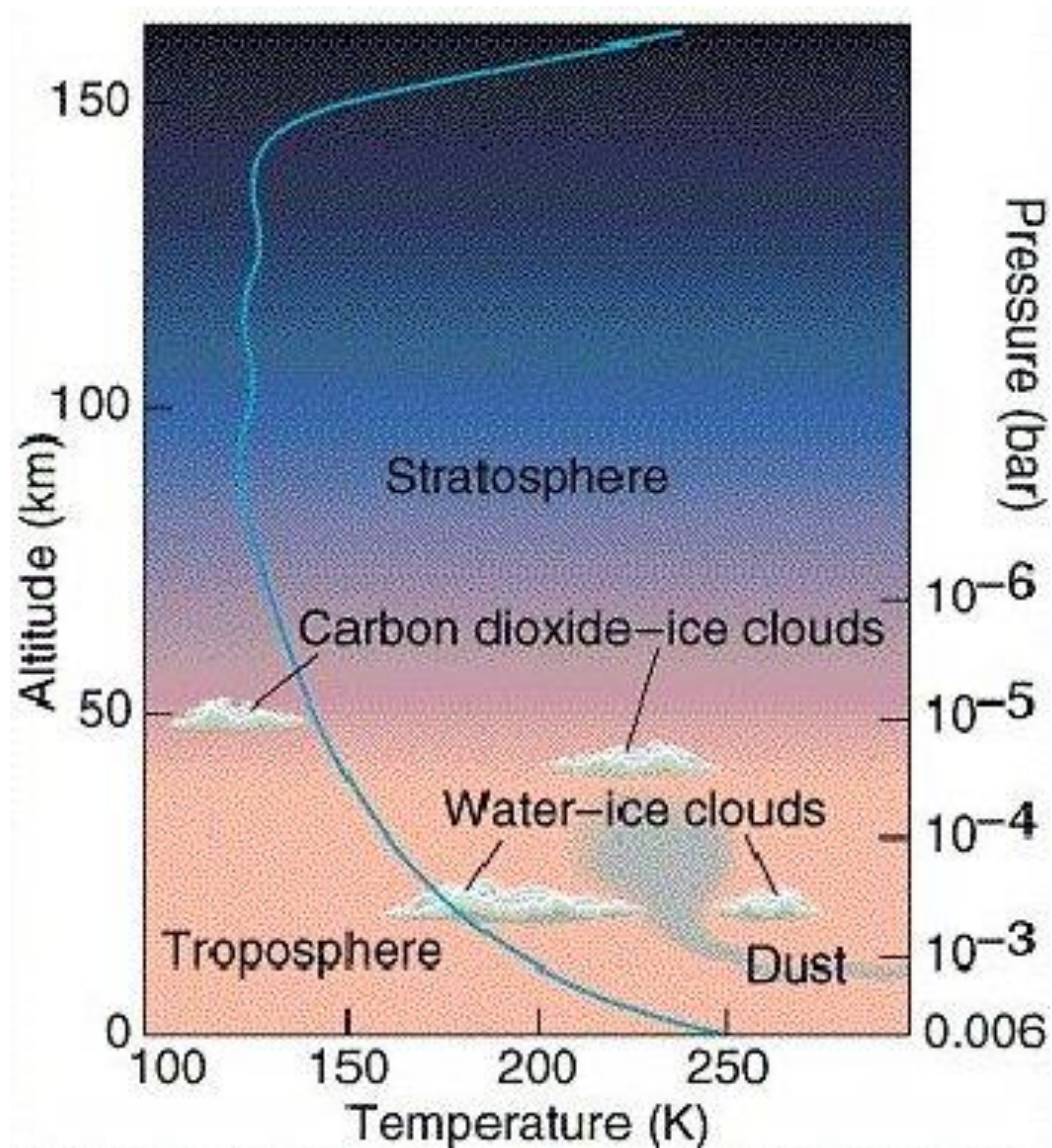
Belgian Institute for Space Aeronomy



Mars

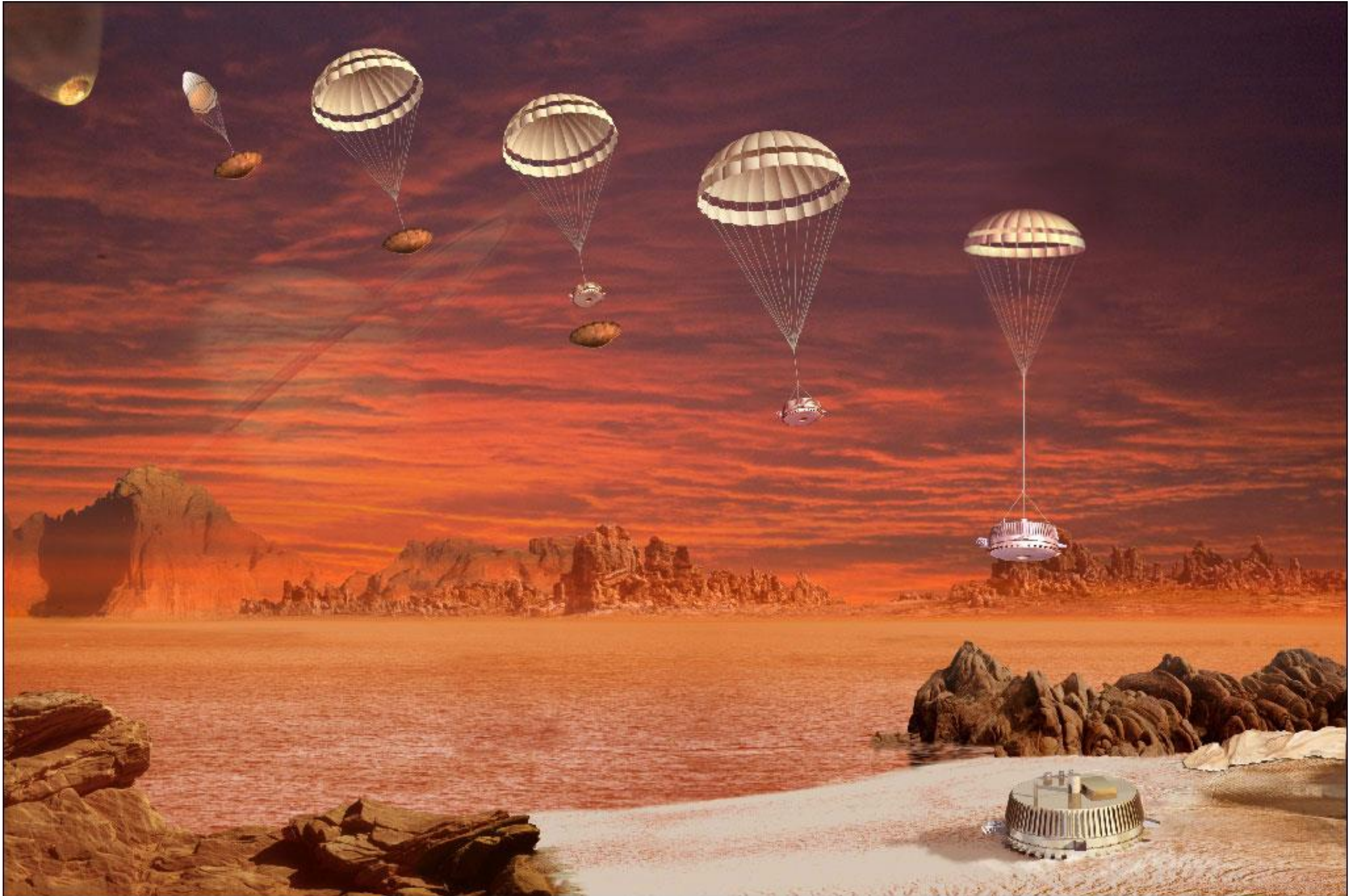
Altitude versus Temperature

<http://astronomy.nmsu.edu/tharriso/ast105/Mars.html>



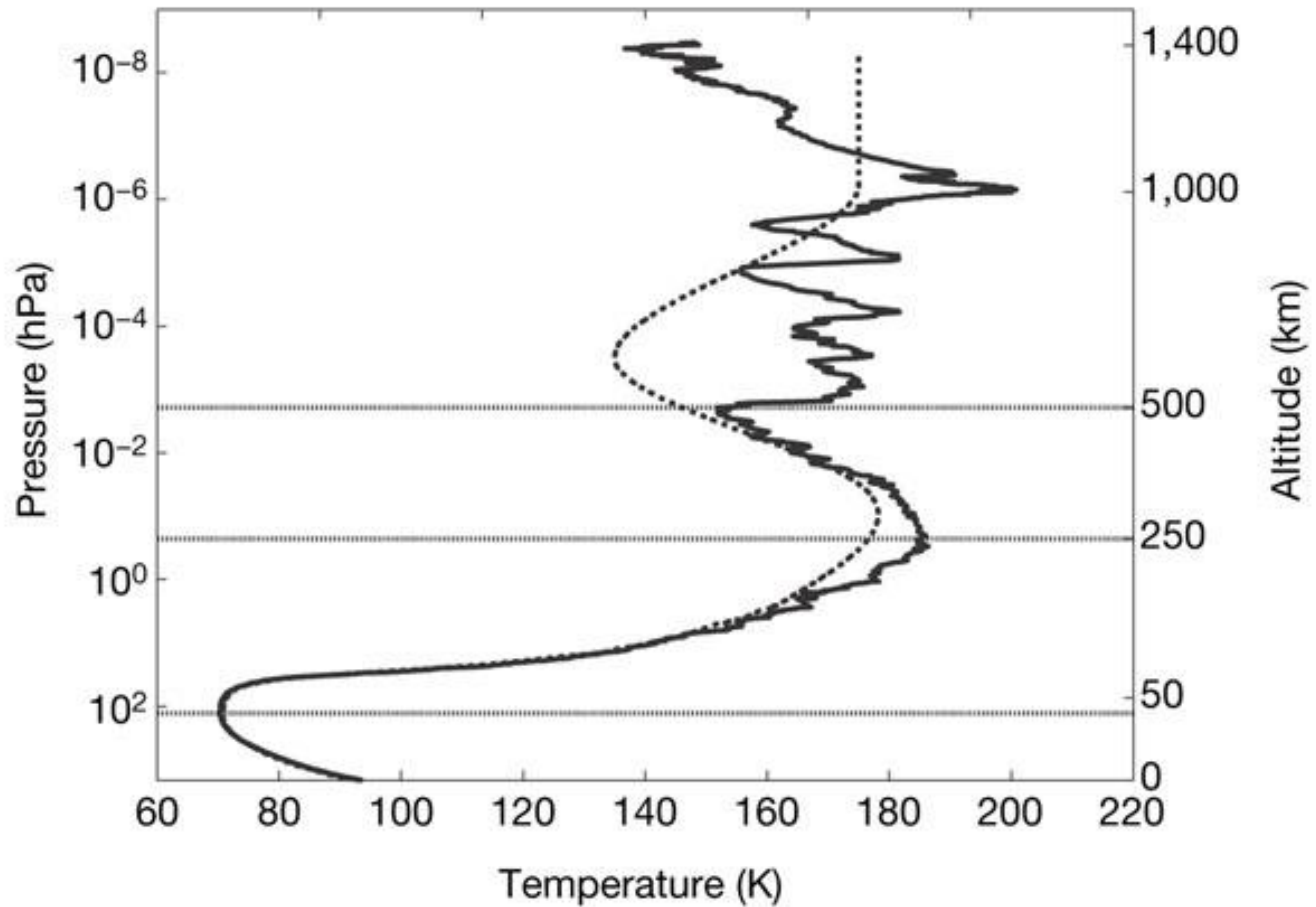
Huygen's descent to the surface of Titan

14 January 2005



Titan

Altitude versus Temperature



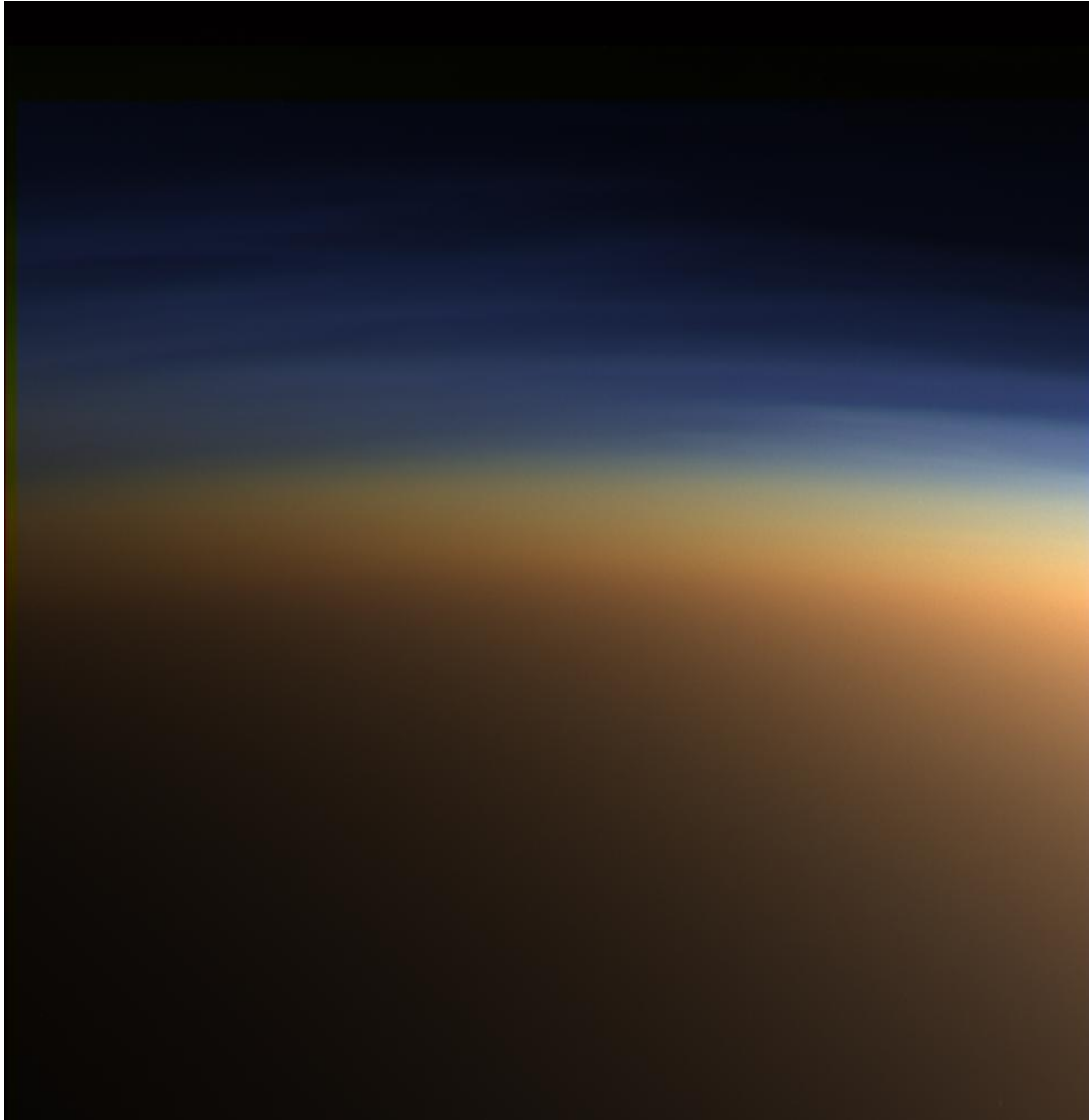
Stratospheric Temperature Increase in Titan's Atmosphere

Heating due to solar ultraviolet radiation
photo-dissociation of CH_4 .

Chemical reaction produce hydrocarbon-nitrile aerosols known as tholins, for example $(\text{C}_{11}\text{N}_4\text{H}_{14})_n$.

Tholins generate the orange color of Titan's atmosphere

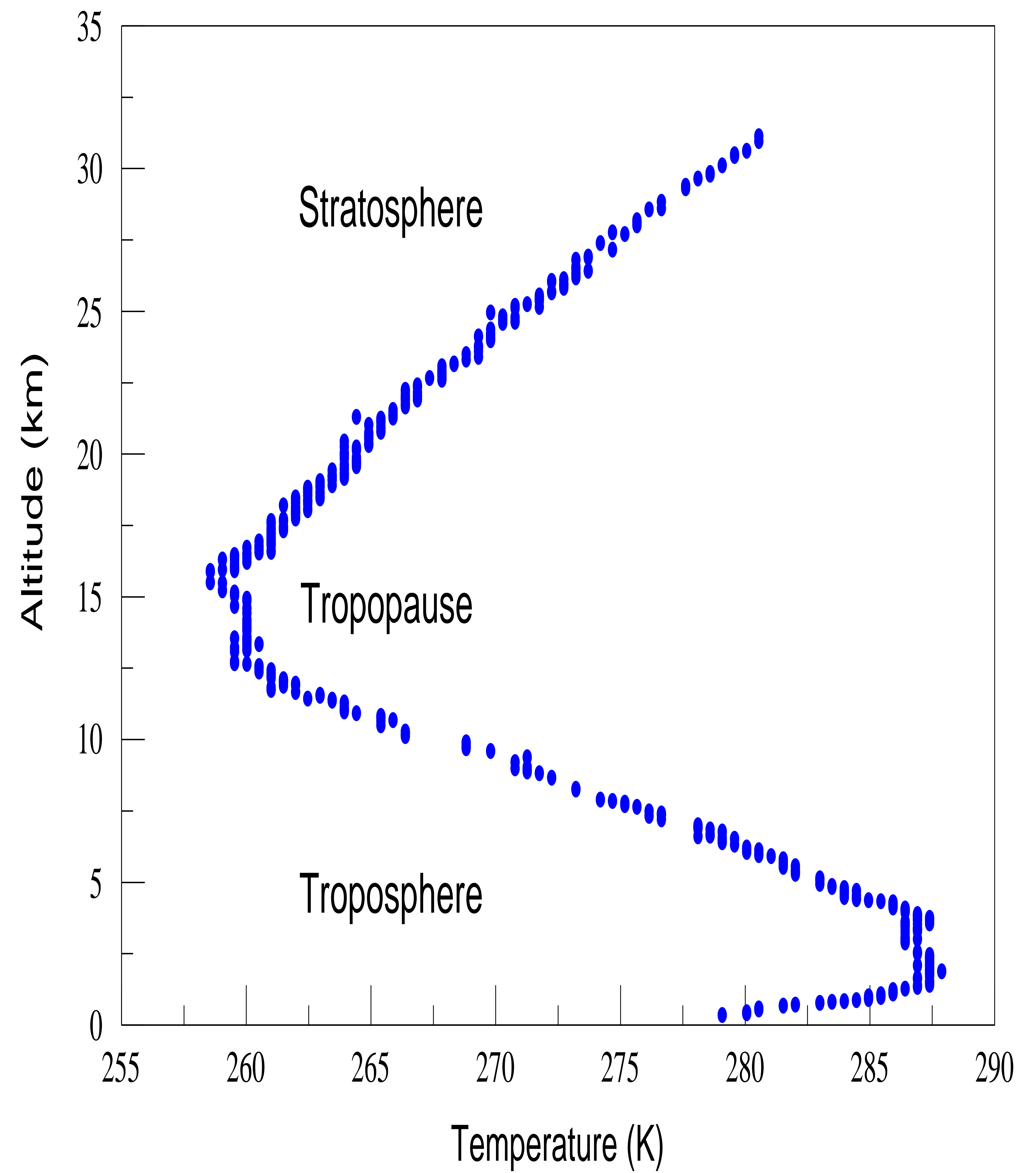
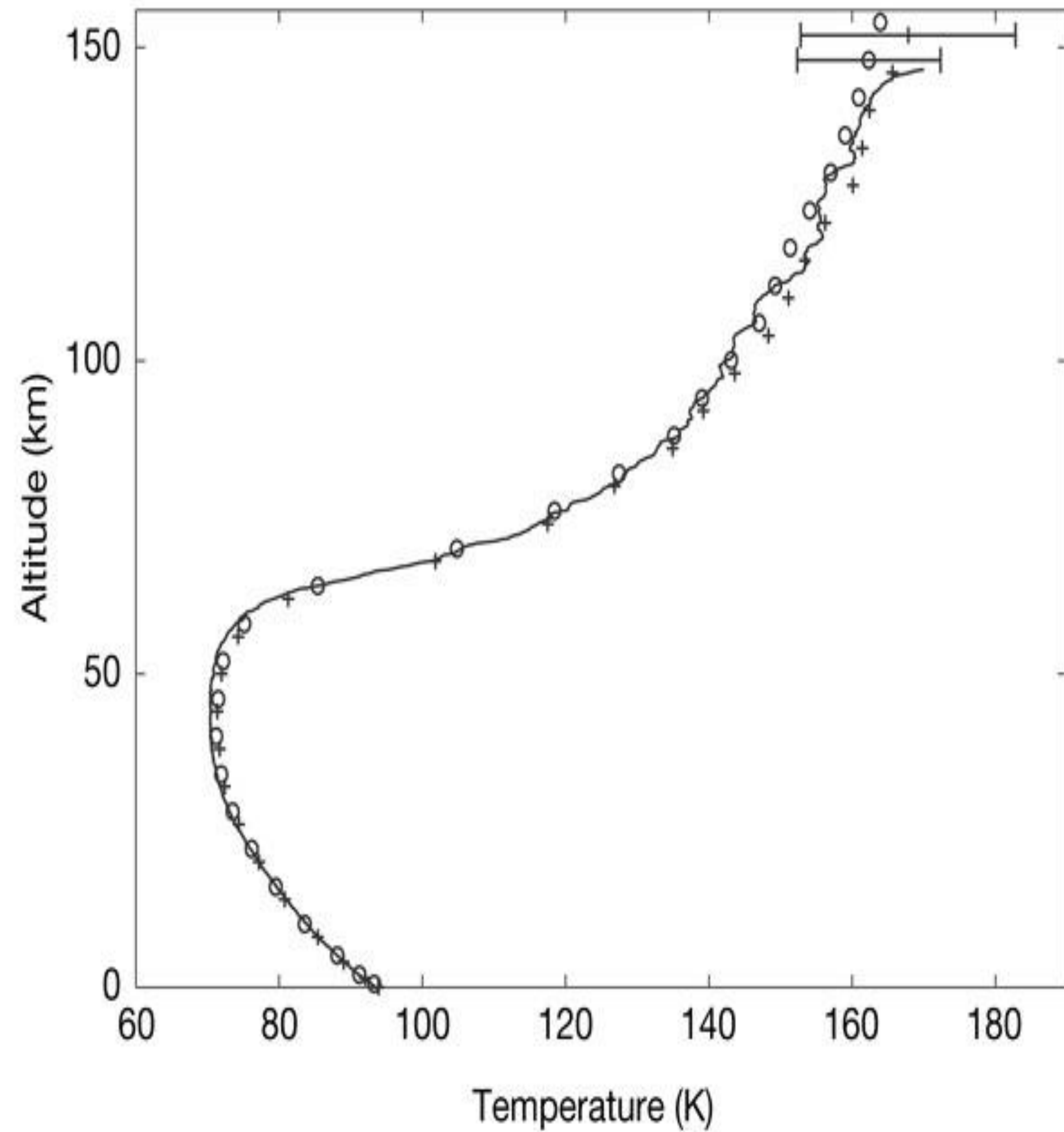
Titan's Atmosphere



Titan's Surface



Temperature profiles for the lower atmospheres of Titan and the Earth



$ELR_{\text{Titan}} < 0.85 \text{ K km}^{-1}$ near the surface

$g_{\text{Titan}} = 1.36 \text{ m s}^{-2}$

$C_p \text{ Titan} = 1044 \text{ J Kg}^{-1} \text{ K}^{-1}$

$DALR_{\text{Titan}} = 1.30 \text{ K km}^{-1}$

SALR =? Methane saturation not water

Lab Activities

- Temperature measurement and calibration
- Temperature structure of the Earth's atmosphere
 - Lapse rates
 - Stability
- Temperature structure of Titan's atmosphere
 - Lapse rates
- Physical and chemical processes

References

- Briggs, D., Smithson, P., Addison, K., and Atkinson, K, ***Fundamentals of the Physical Environment, Second edition***, Routledge Publishing, New York, NY, 1997.
- M. Fulchignoni, et al. *Nature* **438**, 785-791 (8 December 2005) ***In situ measurements of the physical characteristics of Titan's environment***
- **European Space Agency: Titan images**
- http://pds-atmospheres.nmsu.edu/education_and_outreach/encyclopedia/adiabatic_lapse_rate.htm The Planetary Atmosphere Data Node [cited 23 May 2011]
- Lucy-Ann McFadden, Paul R. Weissman and Torrence V. Johnson, Editors, ***Encyclopedia of the Solar System***, Academic Press, Boston, MA, 2007, Chapters 7, 9, 15, and 25
- NSF CCLI Balloon Workshop, Taylor University, 29 – 30 July, 2009, Reference documents and references therein